

# Identifying Loci for Ocean Forecast Model Error from Ensembles of Surface Winds based on QuikSCAT and COAMPS

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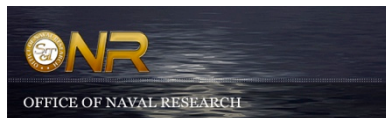
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Milliff, R.F., A. Bonazzi, C.K. Wikle, N. Pinardi and L.M. Berliner, 2011: Ocean Ensemble Forecasting, Part I: Ensemble Mediterranean Winds from a Bayesian Hierarchical Model., *Quarterly Journal of the Royal Meteorological Society*, **137**, Part B, 858-878, doi: 10.1002/qj.767.

Moore, A.M., H.G. Arango, G. Broquet, B.S. Powell, J. Zavala-Garay and A.T. Weaver, 2011: The Regional Ocean Modeling System (ROMS) 4-dimensional variational data assimilation systems. Part I: System overview and formulation. *Progress in Oceanography*, **91**, 34-49.

Moore, A.M., H.G. Arango, G. Broquet, C. Edwards, M. Veneziani, B. Powell, D. Foley, J. Doyle, D. Costa and P. Robinson, 2011: The Regional Ocean Modeling System (ROMS) 4-dimensional variational data assimilation systems. Part II: Performance and application to the California Current System. *Progress in Oceanography*, **91**, 50-73.

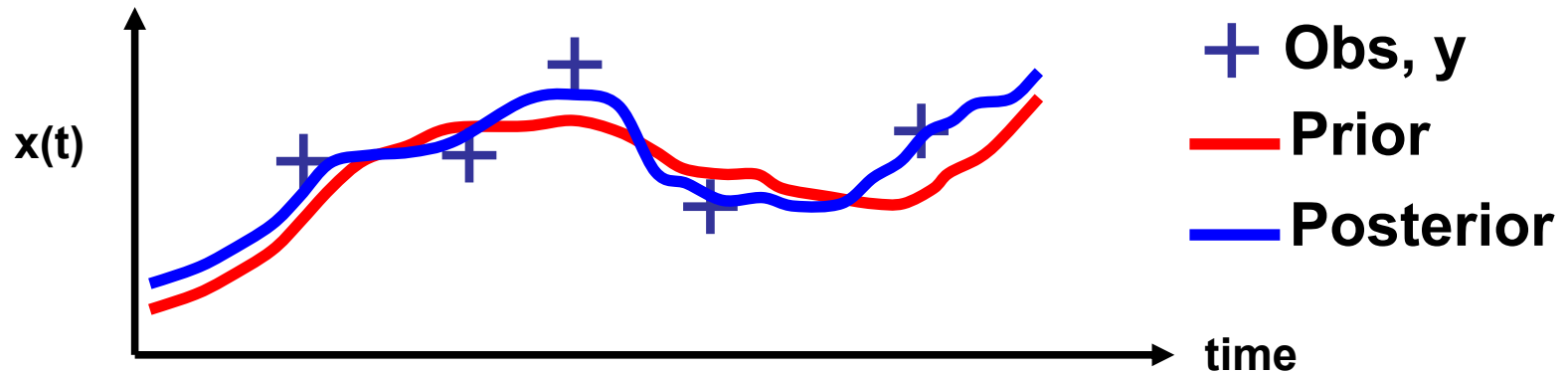
Moore, A.M., H.G. Arango, G. Broquet, C. Edwards, M. Veneziani, B. Powell, D. Foley, J. Doyle, D. Costa and P. Robinson, 2011: The Regional Ocean Modeling System (ROMS) 4-dimensional variational data assimilation systems. Part III: Observation impact and observation sensitivity in the California Current System. *Progress Oceanography*, **91**, 74-94.



## Ocean Data Assimilation (DA); Four-Dimensional Variational (4DVar) Method (in words):

Goal of DA, given state vector  $\mathbf{x}(t-1)$ , is to adjust model trajectory by incorporating observations  $\mathbf{y}(t)$ , such that new estimate of  $\mathbf{x}(t)$  is “optimal”.

So,  $\mathbf{x}_a = \mathbf{x}_b + \mathbf{Kd}$  where  $\mathbf{x}_b$  is the background state (or prior),  $\mathbf{x}_a$  is the new state (or posterior) and  $\mathbf{d} = (\mathbf{y} - \mathbf{Hx}_b)$  are the innovations due to observations.



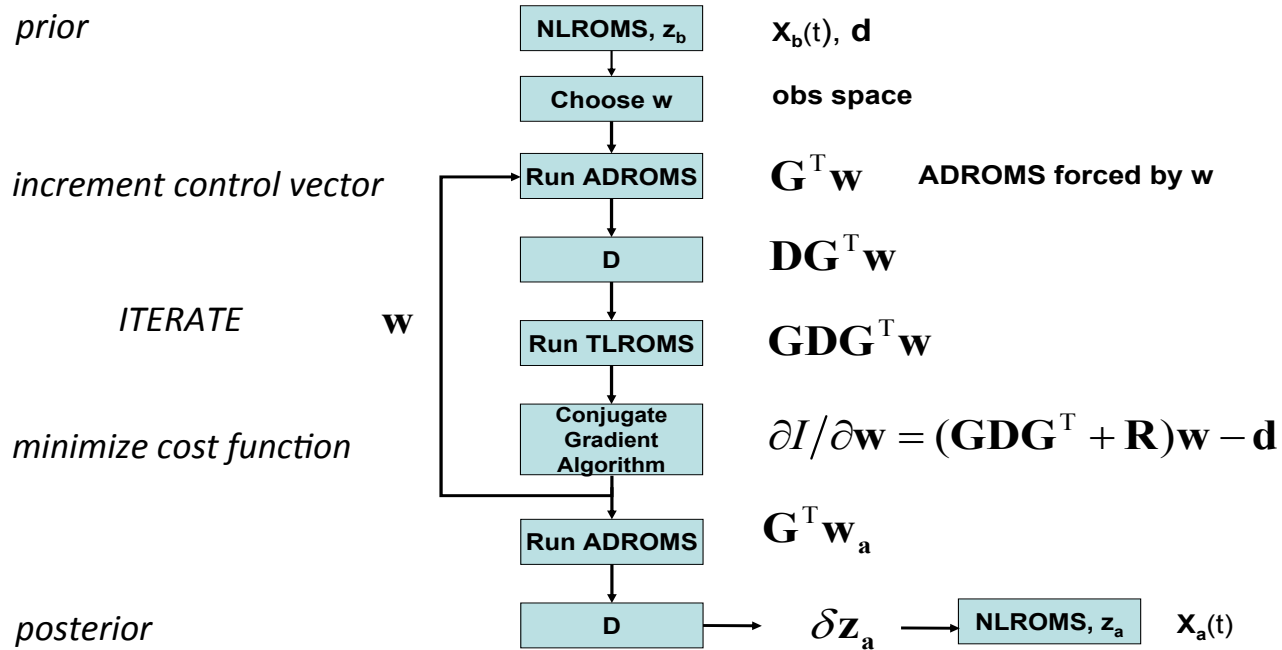
4DVar is a DA method wherein  $\mathbf{Kd}$  is expressed as a control vector of increments:

$\delta\mathbf{z} = [\delta\mathbf{x}(0), \delta\mathbf{f}(t_i), \delta\mathbf{b}(t_i), \delta\eta(t_i)]^T$  for increments in initial conditions  $\delta\mathbf{x}(0)$ , forcing  $\delta\mathbf{f}(t_i)$ , boundary conditions  $\delta\mathbf{b}(t_i)$  and model error  $\delta\eta(t_i)$ . The model error increments are very hard to estimate and are neglected in “strong constraint” 4DVar.

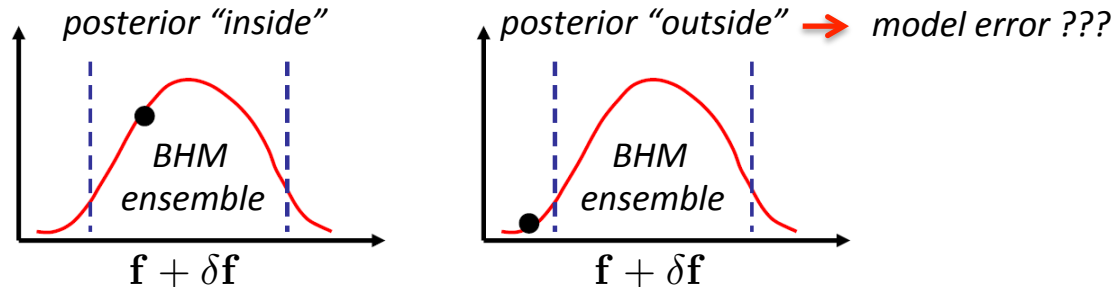
The posterior is “maximized” by minimizing a cost function. Minimization is achieved iteratively by incrementing the control vector. We will compare increments in the surface wind stress part of the control vector  $\delta\mathbf{f}$  with ensemble winds from a

Bayesian Hierarchical Model (BHM)

**Hypothesis:** In a strong constraint 4DVar, places where the wind stress control vector iterations are driven outside a BHM wind ensemble are candidate loci for model error.



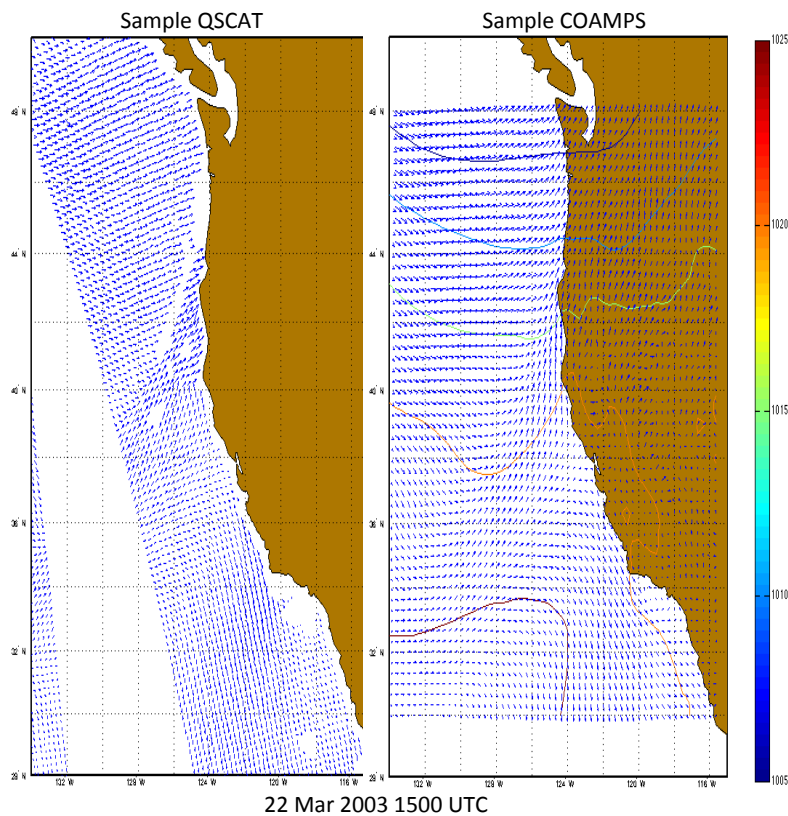
**Schematic:**



# Ensemble Winds from a Bayesian Hierarchical Model (BHM)

$$[X, \vartheta_p, \vartheta_d | D] \propto [D | X, \vartheta_d] [X | \vartheta_p] [\vartheta_p] [\vartheta_d]$$

*posterior*
*likelihood*
*prior*
*parameters*



**Data Stage**  
**(likelihood)**

## Rayleigh Friction Equations

$$\frac{\partial u}{\partial t} - f v = -\frac{1}{\rho_o} \frac{\partial p}{\partial x} - \gamma u$$

$$\frac{\partial v}{\partial t} + f u = -\frac{1}{\rho_o} \frac{\partial p}{\partial y} - \gamma v$$

## EOF expansion SLP anomaly

$$p(x, y, t) = \mu + \sum_{k=1}^m a_k(t) \phi_k(x, y).$$

## Geostrophic-ageostrophic approximation

$$u = -\frac{f}{\rho_o(f^2 + \gamma^2)} \frac{\partial p}{\partial y} - \frac{\gamma}{\rho_o(f^2 + \gamma^2)} \frac{\partial p}{\partial x}$$

$$v = \frac{f}{\rho_o(f^2 + \gamma^2)} \frac{\partial p}{\partial x} - \frac{\gamma}{\rho_o(f^2 + \gamma^2)} \frac{\partial p}{\partial y}$$

## Stochastic Model

$$\mathbf{U}_t = a_{1,1} \mathbf{D}_y \mathbf{P}_t + a_{1,2} \mathbf{D}_x \mathbf{P}_t + \epsilon_{u,t};$$

$$\mathbf{V}_t = b_{1,1} \mathbf{D}_x \mathbf{P}_t + b_{1,2} \mathbf{D}_y \mathbf{P}_t + \epsilon_{v,t};$$

$$P_t(x, y) = \mu + \sum_{k=1}^N \alpha_{k,t} \phi_k(x, y).$$

## Random variables, uncertainty model

$$a_{1,1} \sim N\left(-\frac{f}{\rho_o(f^2 + \gamma^2)}, \sigma_{a_{11}}^2\right)$$

$$a_{1,2} \sim N\left(-\frac{\gamma}{\rho_o(f^2 + \gamma^2)}, \sigma_{a_{12}}^2\right)$$

$$b_{1,1} \sim N\left(\frac{f}{\rho_o(f^2 + \gamma^2)}, \sigma_{b_{11}}^2\right)$$

$$b_{1,2} \sim N\left(-\frac{\gamma}{\rho_o(f^2 + \gamma^2)}, \sigma_{b_{12}}^2\right)$$

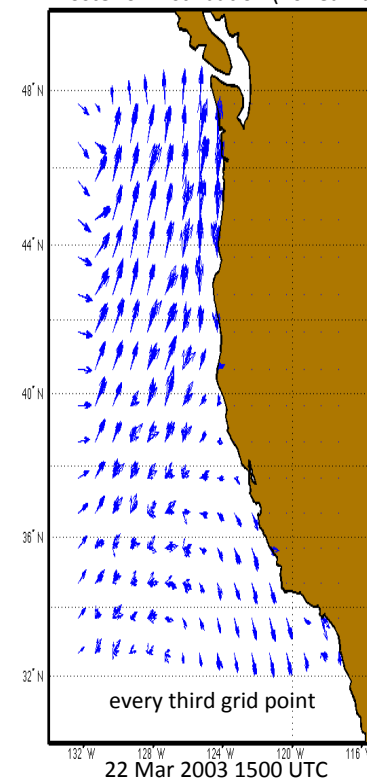
$$\alpha_{k,t} \sim N(0, \sigma_P^2).$$

$$\epsilon_{u,t} = \sum_{k=1}^{n_\beta} \mathbf{W}_k \beta_{k,t}^u + \tilde{\epsilon}_{u,t}$$

$$\tilde{\epsilon}_{u,t} \sim N(\mathbf{0}, \sigma_u^2 \mathbf{I}).$$

**Process Model Stage**  
**(prior)**

## BHM Posterior Distribution (10 realizations)

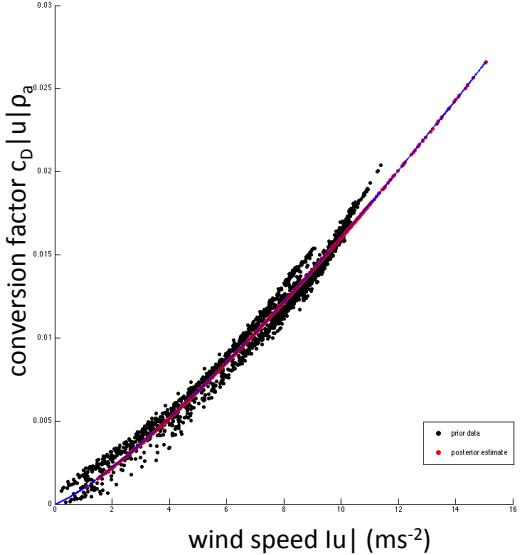
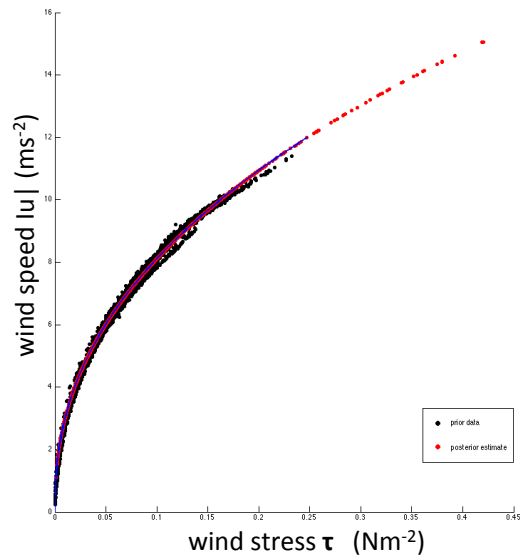


**Posterior**

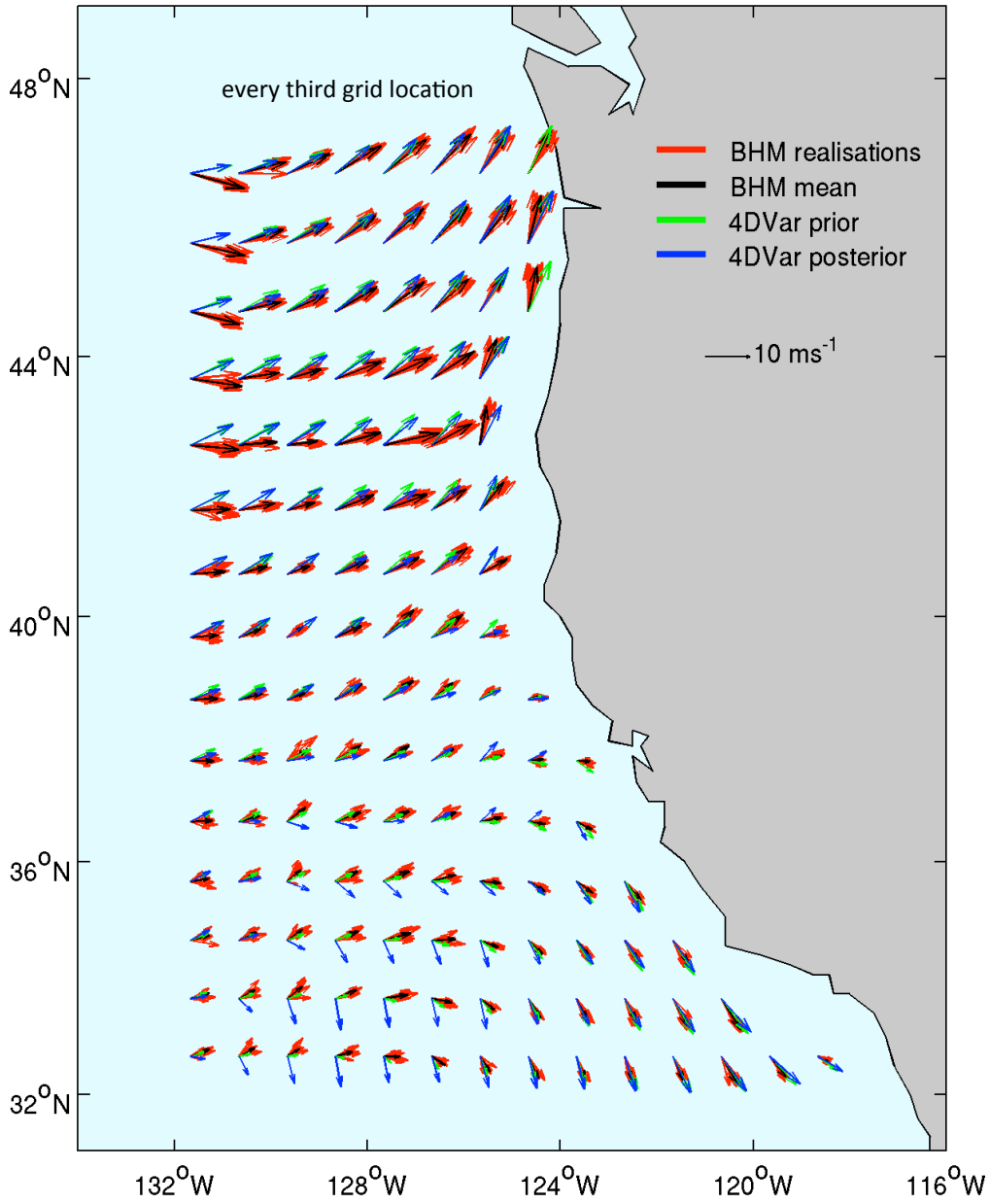
# Comparing 4dVar Increments with BHM

## Estimate u,v from $\tau$ every day

1. find  $|u|$  as function of  $\tau$
2. find  $c_D|u|\rho_a$  as function of  $|u|$
3. estimate u,v from posterior  $\tau_{x,y} / c_D|u|\rho_a$



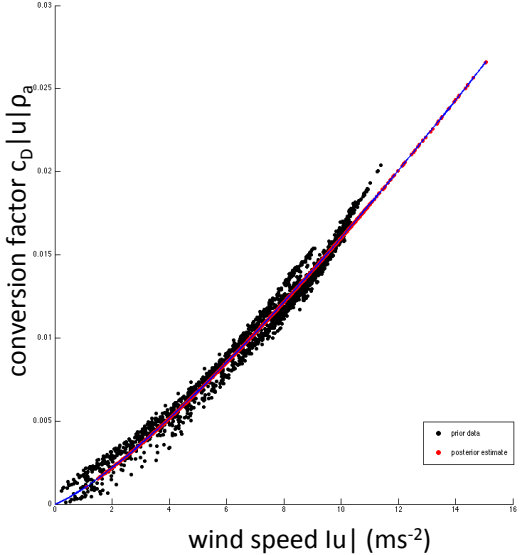
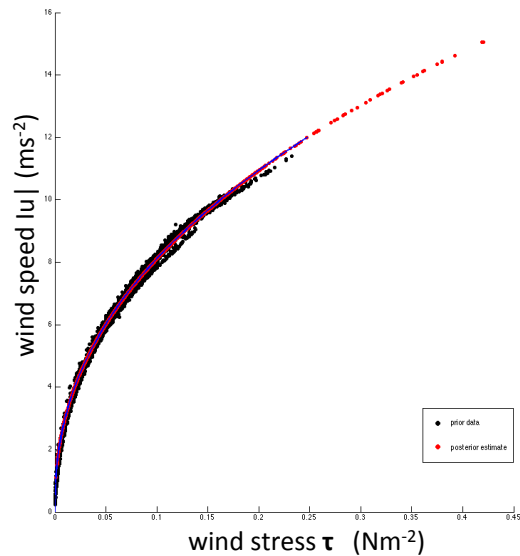
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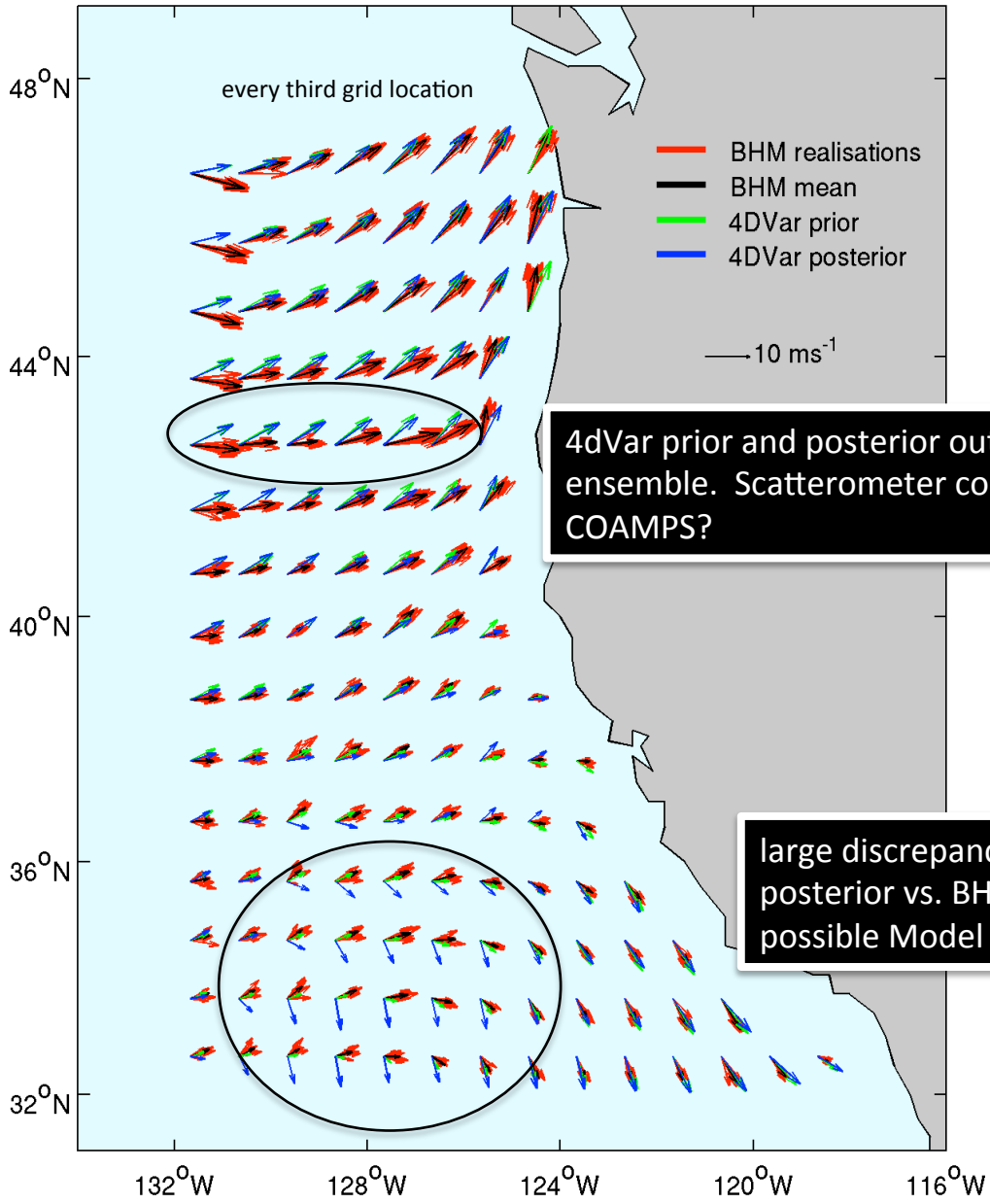
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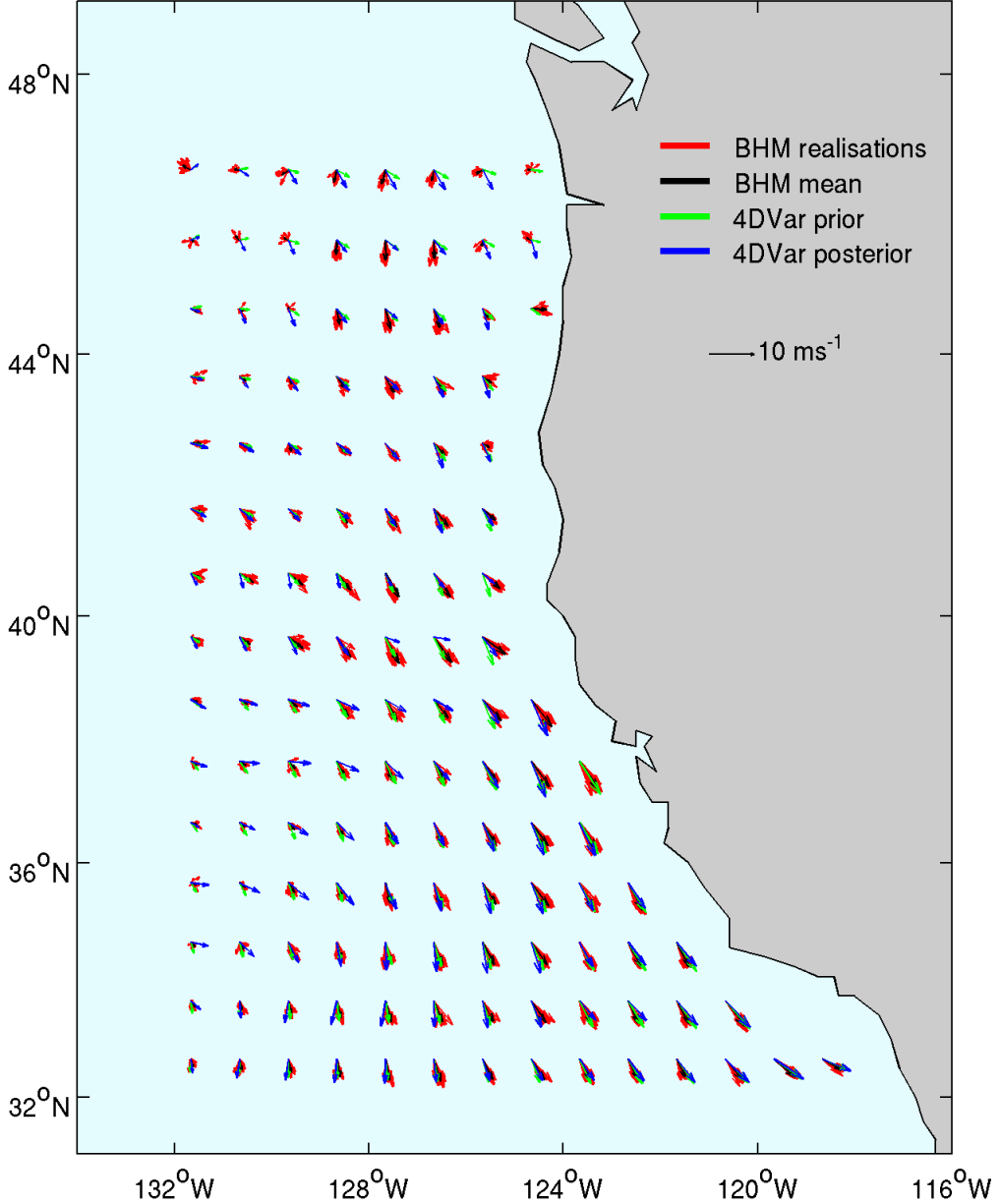


4dVar prior and posterior outside BHM ensemble. Scatterometer correction to COAMPS?

large discrepancies, 4dVar posterior vs. BHM ensemble; possible Model Error

Model Error as a dynamical function of space, time

01-Mar-2003: jdate 12700



## Summary

- ❑ Ocean forecast model error estimates are important for assessing forecast accuracies and improving model and observation array designs (strong vs. weak constraint in 4dVar).
- ❑ Model error is hard to identify.
- ❑ Precisely specified error properties (uncertainty) in QuikSCAT winds have been used to help identify regions of probable model error in a state-of-the-art ocean forecast system.
- ❑ The probability distributions from a surface wind BHM (based on QuikSCAT and COAMPS) have quantitative value (see also Pinardi et al., 2011).

## Work in Progress

- ✧ Wind stress distributions can be summarized directly from surface wind BHM posterior
- ✧ Other forcing control variable fluxes require BHMs as well (e.g. heat, fresh water)
- ✧ Model error dynamics can be modeled given data stages of the kinds suggested here.



**EXTRAS**

# Temporal Autocorrelation Comparison: BHM vs. COAMPS (all of 2003)

AR-1 model for temporal autocorrelation:  $U(t; x, y) = r_1(x, y)U(t - 1; x, y) + \epsilon(t - 1)$

$r_1(x, y)$

